

# Antarctic Science Bursary Report

## Provenance of Aeolian Dust Deposited in McMurdo Sound, SW Ross Sea

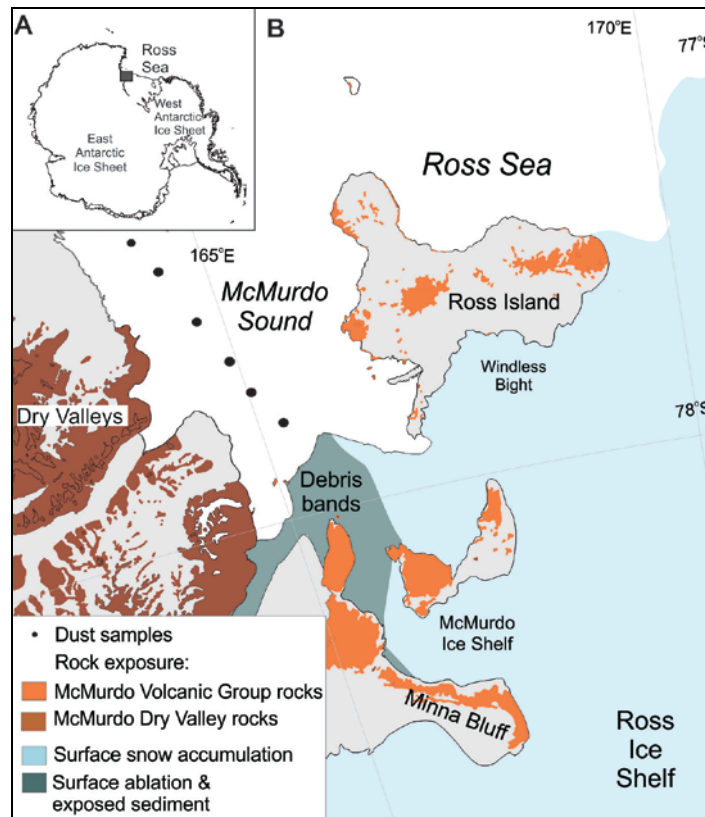
VHL Winton  
PhD candidate  
Curtin University, Perth, Australia

### Project description

Aeolian dust is a source of soluble iron (Fe), which is the limiting nutrient required for phytoplankton growth in high nutrient low chlorophyll (HNLC) regions of the world's ocean including the Southern Ocean.<sup>1-5</sup> This 'iron fertilisation' of the ocean results in vast phytoplankton blooms that alter the food web and have the potential to export particulate carbon to the deep ocean.<sup>6</sup> Each summer the Fe-limited waters of McMurdo Sound, in the south-western (SW) Ross Sea, bloom with phytoplankton.<sup>7-8</sup> This phenomenon is thought to be stimulated by the addition of soluble Fe. An important source of soluble Fe is aeolian dust from local and global sources.<sup>e.g. 9</sup> The solubility of Fe is influenced by several parameters including dust flux, particle size distribution and provenance.<sup>e.g. 10-13</sup> Aspects of this biogeochemical cycle have been addressed for the first time by analysing the physical and geochemical properties of the aeolian dust accumulating on sea ice, deposited and trapped in coastal snow and ice in the McMurdo Sound region<sup>14</sup>.

An important control on Fe solubility is dust provenance.<sup>e.g. 9-12</sup> The geochemical fingerprint of remote dust deposited on the Antarctic plateau shows it originates from arid regions in Patagonia and Australia,<sup>15-18</sup> but contribution from exposed Antarctic sources is likely for sites located on the margin of the ice sheet.<sup>19</sup> Global 'background dust' is characterized by its fine grain size (mode  $<5 \mu\text{m}$ ) and modern mass accumulation rates of  $0.001\text{-}0.02 \text{ g m}^{-2} \text{ yr}^{-1}$ .<sup>20-24</sup> In contrast, the debris bands on the McMurdo Ice Shelf represent the most significant local dust source in Antarctica and contributes  $\sim 1.5 \text{ g m}^{-2} \text{ yr}^{-1}$  of aeolian dust to SW Ross Sea (Fig. 1).<sup>14</sup>

There are two local potential source areas (PSAs) in McMurdo Sound (a western source: McMurdo Dry Valleys and a southern source: McMurdo Volcanic Group) each with a distinctive geochemical fingerprint. The combination of different isotopic signatures ( $^{87}\text{Sr}/^{86}\text{Sr}$  and  $^{143}\text{Nd}/^{144}\text{Nd}$ ) is extremely useful in tracing PSAs because different geographic provenances can be discriminated by variations in radiogenic isotopes of mantle derived (basaltic rocks, tephra and soils derived from them, weathered and eroded mafic particles) or crustal (soils, sediments) derived sediments.<sup>25</sup> This bursary was used to determine the Sr and Nd isotopic signature of McMurdo Sound dust and compare the results with the isotopic signature of local PSAs.



**Fig. 1:** A) Map of Antarctica showing location of McMurdo Sound. B) Map of McMurdo Sound showing the north-south transect of dust samples on the sea ice and potential source areas.

## Method

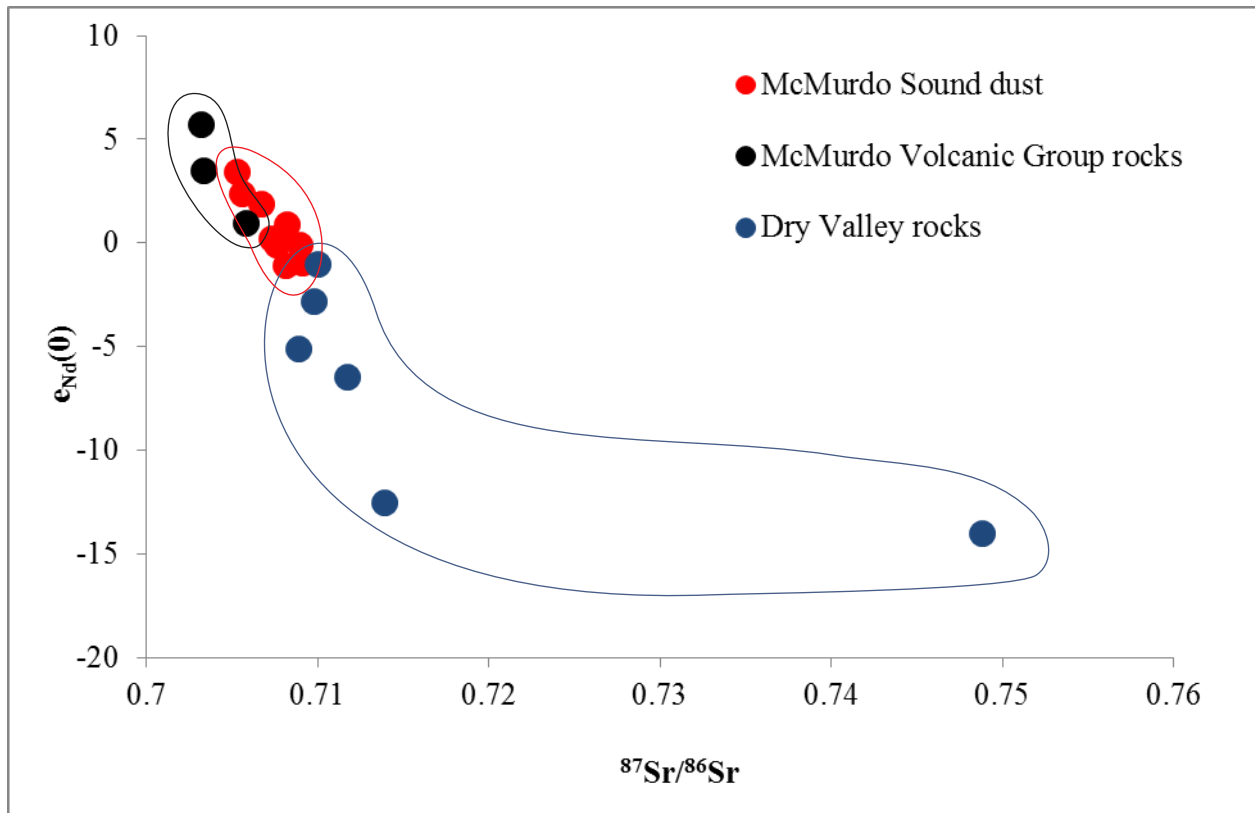
Dust sample digestion and separation of Nd and Sr using ion exchange chromatography<sup>26</sup> was carried out at the Laboratory for Isotope Geology at the Swedish Museum of Natural History, Stockholm under clean conditions in a class 100 laminar flow hood. The samples were spiked with <sup>150</sup>Nd and <sup>84</sup>Sr for isotope dilution determination of the concentrations and analysed on a Thermo Scientific TRITON thermal ionisation mass-spectrometer (TIMS). Reproducibility of the <sup>143</sup>Nd/<sup>144</sup>Nd and <sup>87</sup>Sr/<sup>86</sup>Sr ratios is 50 ppm and 40 ppm respectively. The total column yield for the Nd and Sr separation exceed 95 %. The total blank concentrations, including dissolution, chemical separation and mass spectrometry, was <10 pg for Nd and <80 pg for Sr.

## Results and discussion

The Sr and Nd isotopic ratios of aeolian dust in McMurdo Sound are reported in Fig. 2 and compared with PSA samples from Delmonte et al.<sup>19, 27</sup> The large particle size and high dust flux suggest transport distance is limited and therefore only local PSAs are considered. Sedimentological evidence points to the debris bands as the most important source of dust in the McMurdo Sound region.<sup>14</sup> The additional isotopic analyses confirm dust accumulating on the sea ice is locally-derived. This data is used to investigate the relative contribution of the two dust sources in McMurdo Sound by considering the samples as a mixture between two end members: McMurdo Dry Valleys and McMurdo Volcanic Group (Figs. 1 and 2). Overall, the Sr isotopic ratios for McMurdo Sound PSAs range between  $0.703 < ^{87}\text{Sr}/^{86}\text{Sr} < 0.714$  while  $\epsilon\text{Nd}(0)$  ranges between  $(6 < \epsilon\text{Nd}(0) > -13)$ . Assuming these two end members, it is observed that the dust samples are included within the mixing field: the isotopic signature  $(0.705331 < ^{87}\text{Sr}/^{86}\text{Sr} < 0.7091465$  and  $-1.1 < \epsilon\text{Nd}(0) > 3.45)$  lies on the mixing line between the two isotopically distinct end members. Within the narrow range of isotopic ratios of McMurdo Sound dust, there is a small change in the <sup>87</sup>Sr/<sup>86</sup>Sr and <sup>143</sup>Nd/<sup>144</sup>Nd composition of the dust along the south-north transect (Fig. 1). As the dust is transported north from the debris bands the composition moves along the mixing line from the

McMurdo Volcanic Group towards the Dry Valleys. This gradient in mixing is consistent with the exponential decrease in dust flux downwind from the debris bands and the local meteorology.

The narrow range of isotopic ratios has implications for supply of Fe to the SW Ross Sea. A previous study observed that there is little variability in Fe solubility<sup>14</sup> along the north-south transect in Fig. 1 and the homogenous composition of the dust can thus explain this. Given that total Fe content and particle size do not influence Fe solubility<sup>14</sup>, it can be concluded that dust provenance is the major control on Fe solubility in McMurdo Sound. This conclusion is consistent with studies from other regions.<sup>e.g.28</sup>



**Fig. 1: Isotopic composition of McMurdo Sound dust (red circles) compared with PSA samples from Delmonte et al.<sup>20,28</sup>**

### Conclusion and future work

Local dust accumulating on McMurdo Sound sea ice is derived from exposed areas of unconsolidated sediment. The isotopic signature of McMurdo Sound dust lies between two isotopically distinct PSAs in the sound and thus is a mixture of these two sources. The isotopic composition of McMurdo Sound dust falls within a narrow range ( $0.705331 < ^{87}Sr/^{86}Sr < 0.7091465$  and  $-1.1 < \epsilon_{Nd}(0) > 3.45$ ). This homogenous composition can explain the constant Fe solubility of the dust samples along a north-south transect. If this local dust is transported beyond McMurdo Sound and further north into the Ross Sea, Fe supplied from the local dust can be quantified from the known Fe solubility. Samples from Roosevelt Island, on the eastern side of the Ross Ice Shelf, will be analysed for  $^{87}Sr/^{86}Sr$  and  $^{143}Nd/^{144}Nd$  isotopic ratios to determine how far north locally derived dust is transported into the Ross Sea. The broader implications of this research are to determine the influence of the dust source on Fe solubility in McMurdo Sound and to quantify the area in which the dust is potentially bio-available for phytoplankton growth in the Ross Sea, a critical region of carbon sequestration.

### References

1. Martin, J, Gordon. M and Fitzwater. S, 1991. The case for iron. *Limnology Oceanography*, 36, 1793-1802.
2. Sedwick. P, and DiTullio. G, 1997. Regulation of algal blooms in Antarctic shelf waters by the release of iron melting sea ice. *Geophysical Research Letters*, 24, 2515-2518.
3. Lannuzel, D., V. Schoemann, J. de Jong, J.-L. Tison, and L. Chou, 2007. Distribution and biogeochemical behaviour of iron in the East Antarctic sea ice, *Marine Chemistry*, 106(1-2), 18-32.
4. Tagliabue, A., L. Bopp, and O. Aumont, 2009. Evaluating the importance of atmospheric and sedimentary iron sources to Southern Ocean biogeochemistry, *Geophys. Res. Lett.*, 36(13), L13601.
5. Boyd, P. W., and M. J. Ellwood, 2010. The biogeochemical cycle of iron in the ocean, *Nature Geosci*, 3(10), 675-682.
6. Smetacek, V., et al., 2012. Deep carbon export from a Southern Ocean iron-fertilized diatom bloom, *Nature*, 487(7407), 313-319.
- 7 Arrigo. K and, van Dijken. G, 2004. Annual changes in sea-ice, chlorophyll a, and primary production in the Ross Sea, Antarctica. *Deep-Sea Research II*, 51, 117-138.
8. Fung. I, Meyn. S, Tegan. I, Doney. S, John. J, Bishop. J, 2000. Iron supply and demand in the upper ocean. *Global Biogeochemical Cycles*, 14, 281-295.
9. Baker, A, Jickells, T, Witt, M., Linge, K, 2006. Trends in the solubility of iron, aluminium, manganese, and phosphorus in aerosol collected over the Atlantic Ocean. *Marine Chemistry*. 98, 43–58.
- 10 Raiswell. R, Benning. L, Tranter. M and Tulaczyk. S, 2008. Bioavailable iron in the Southern Ocean: the significance of the iceberg conveyor belt. *Geochemical Transactions*, 9:7, doi:10.1186/1467-4866-9-7.
11. Aguilar-Islas. M, Wu. J, Rember. R, Johansen. A, and Shank. L. 2009. Dissolution of aerosol-derived iron in seawater: Leach solution chemistry, aerosol type, and colloidal iron fraction. *Marine Chemistry*, 120, 25-33.
12. Schroth. A, Sholkovitz. E and Bostick. B, 2009. Iron solubility driven by speciation in dust sources to the ocean. *Nature Geoscience*, 2, 337-370.
13. Buck. C, Landing. W, Resing. J, 2010. Particle size and aerosol iron solubility: A high-resolution analysis of Atlantic aerosols. *Marine Chemistry*, 120, 14-24.
14. Winton. V.H.L, Dunbar. G.B, Bertler, N.A.N, Millet. M-A, Delmonte. B, Atkins. C.B, Andersson. P, in review. The contribution of locally-derived aeolian sand and dust to Fe fertilisation of phytoplankton blooms in southwest Ross Sea. *Global Biogeochemical Cycles*.
15. Basile. I, Grousset. F, Revel. M, Petit. J, Biscaye. P, Barkov. N, 1997. Patagonian origin of glacial dust deposited in East Antarctica (Vostok and Dome C) during glacial stages 2, 4 and 6. *Earth and Planetary Science Letters*, 146, 573-589.
16. De Deckker. P, Norman. N, Goodwin. I, Wain. A, Gingele. F, 2010. Lead isotopic evidence for an Australian source of aeolian dust to Antarctica at times over the last 170,000 years. *Palaeogeography, Palaeoclimatology, Palaeoecology*, 285, 205–223.
17. Delmonte. B, Basile-Drelsh. I, Petit. J, Maggi. V, Revel-Rolland. M, Michard. A, Jogoutz. E and Grousset. F, 2004. Comparing the EPICA and Vostok dust records during the last 220,000 years: stratigraphical correlation and provenance in glacial periods. *Earth Science Reviews*, 66, 63-87.
18. Vallelonga, P., et al., 2010. Lead isotopic compositions in the EPICA Dome C ice core and Southern Hemisphere Potential Source Areas, *Quaternary Science Reviews*, 29(1-2), 247-255.
19. Delmonte, B., P. S. Andersson, H. Schöberg, M. Hansson, J. R. Petit, R. Delmas, D. M. Gaiero, V. Maggi, and M. Frezzotti, 2010. Geographic provenance of aeolian dust in East Antarctica during Pleistocene glaciations: preliminary

- results from Talos Dome and comparison with East Antarctic and new Andean ice core data, *Quaternary Science Reviews*, 29(1-2), 256-264.
20. Duce, R. A., et al., 1991. The atmospheric input of trace species to the world ocean, *Global Biogeochem. Cycles*, 5(3), 193-259.
21. Mahowald, N.M., Baker, A.R., Bergametti, G., Brooks, N., Duce, R.A., Jickells, T.D., Kubilay, N., Prospero, J.M., Tegen, I., 2005. Atmospheric global dust cycle and iron inputs to the ocean. *Global Biogeochemical Cycles*, 19, GB4025. doi:10.1029/2044GB002402.
22. Wagener, T., C. Guieu, R. Losno, S. Bonnet, and N. Mahowald, 2008. Revisiting atmospheric dust export to the Southern Hemisphere ocean: Biogeochemical implications, *Global Biogeochem. Cycles*, 22(2), GB2006.
23. Albani, S., B. Delmonte, V. Maggi, C. Baroni, J. R. Petit, B. Stenni, C. Mazzola, and M. Frezzotti, 2012. Interpreting last glacial to Holocene dust changes at Talos Dome (East Antarctica): implications for atmospheric variations from regional to hemispheric scales, *Clim. Past Discuss.*, 8(1), 145-168.
24. Delmonte, B., C. Baroni, P. Andersson, B. Narcisi, M. Salvatore, J. Petit, C. Scarchilli, M. Frezzotti, S. Albani, and V. Maggi, 2013. Modern and Holocene aeolian dust variability from Talos Dome (Northern Victoria Land) to the interior of the Antarctic ice sheet, *Quaternary Science Reviews*, 64, 76-89.
25. Grousset. F and Biscaye. P, 2005. Tracing dust sources and transport patterns using Sr, Ns and Pb isotopes. *Chemical Geology*, 222, 149-167.
26. Delmonte, B., P. S. Andersson, M. Hansson, H. Schöberg, J. R. Petit, I. Basile-Doelsch, and V. Maggi, 2008. Aeolian dust in East Antarctica (EPICA-Dome C and Vostok): Provenance during glacial ages over the last 800 kyr, *Geophys. Res. Lett.*, 35(7), L07703.
27. Delmonte. B, Basile-Drelsh. I, Petit. J, Maggi. V, Revel-Rolland. M, Michard. A, Jogoutz. E and Grousset. F, 2004. Comparing the EPICA and Vostok dust records during the last 220,000 years: stratigraphical correlation and provenance in glacial periods. *Earth Science Reviews*, 66, 63-87.
28. Baker, A. R., and P. L. Croot, 2010. Atmospheric and marine controls on aerosol iron solubility in seawater, *Marine Chemistry*, 120(1-4), 4-13.